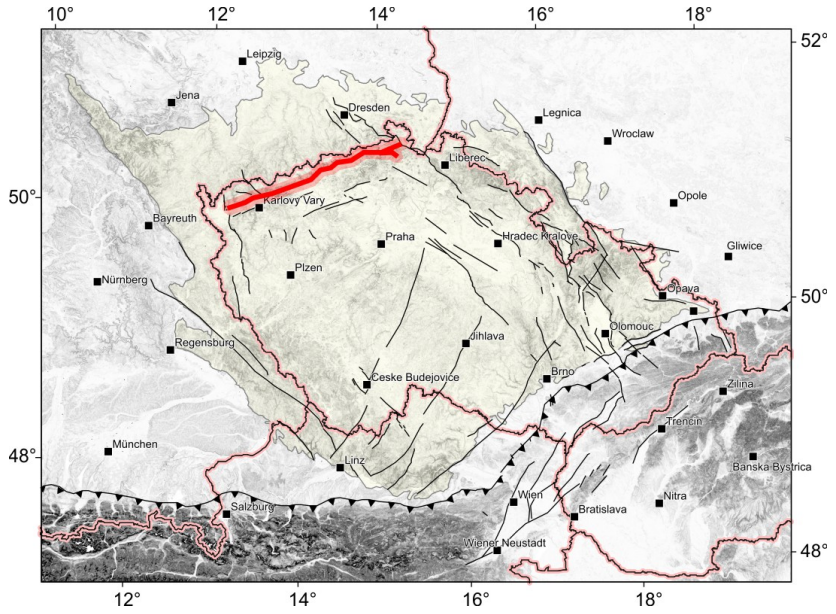
Krušné hory Fault

Structure ID:
Fault Section IDs:
Related terms:

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General description, fault trace and terminology

… Definition: From a terminological point of view, it is appropriate to distinguish the Krušné hory fault (s.s.) from the higher-order fault system, here referred to as the Krušné hory fault, which includes fault structures that follow it in both directions (Fig. 1). In works of a more general nature, this composite system is also referred to as the Krušné hory fault (s.l.).

a) Krušné hory fault s.s. is a fault zone of structurally diverse character, which separates the tectonically higher block of the Krušné hory Mountains on the NNW from the tectonically lower block on the SSE.

The course of the fault is defined in the literature by the municipalities of Horní Žďár in the Sokolovská

basin (Rojík et al. 2010) in the west and Jílová near Děčín in the east (Malkovský 1977, 1979, 1987).

The upper block is represented in the dominant middle part by the exposed crystalline formations of the Krušné Hory Mountains; towards E, probably due to its tectonic inclination, relicts of Upper Cretaceous rocks begin to appear on it, which further change into a continuous cover. In the tectonic lower block, the cover of Upper Cretaceous sediments and tertiary formations of the North Bohemian and Sokolov basins is mostly preserved, which are denuded / eroded to varying degrees in sub-blocks.

b) Krušné Hory fault s.l. is a fault system of a higher order, which usually includes fault structures, which follow on the Krušné Hory Mountains fault s.s. in both directions. They have a similar direction and partly also a structural character. As at the Krušné hory fault itself, there is an expressive decrease in the southern blocks at these subsequent structures at a certain stage of their tectonic development. Especially in works of a more general nature, this complex system is also referred to as the Krušné hory fault. Components and course of the Krušné hory fault s.l .: Towards the west (Fig. 1) to the Krušné hory fault s.s. follows up a system of partial smaller faults in the direction of the WSW - ENE, which participated in the formation of the Cheb basin; however, a clear continuation of the Krušné hory fault into this basin is not evident (Malkovský 1989). In the next continuation, the Krušné hory fault manifests itself as the S boundary of Smrčiny, where it fades in the proximity of Tirschenreuth (Malkovský 1977, Stettner 1974).

Towards the east (Fig. 1) the Krušné hory fault flows smoothly into the Děčín fault and further east into the Česká Kamenice fault field - in both cases, those are wide systems of mostly sub-parallel faults of mainly the E - W direction. These faults are manifested by the violation of sediments of the Czech Cretaceous basin, while the area of conservation of Tertiary sediments in the tectonic lower crust ends towards V in the vicinity of Ústí nad Labem. In the vicinity of Česká Kamenice, this system is followed by a directionally different zone of faults, referred to as the Doubice fault field. This system of faults in the NNE - SSW direction takes over the function of the boundary between the higher NNW to NW circle and the lower SW circle; both shrubs are represented in subsurface parts by Upper Cretaceous sediments of different stratigraphic levels. The faults of the Doubice field end at the Lusatian fault in the section between the municipalities of Kyjov and Brtníky. Their possible continuation through the Lusatian fault to the NE is not traceable in the rocks of the Lusatian pluton (Malkovský 1977, 1979).

The length of the Krušné hory fault system s.l. exceeds 200 km (Malkovský 1977).

The Krušné Hory Mountains fault as a building block of the Eger Graben

The Krušné Hory Mountains fault is undoubtedly one of the basic building blocks of the Eger Graben, because it forms its NW edge fault and its course consistently correspond to the geometry and deep structure of this higher-order volcano-tectonic structure (Fig. 2). According to the basic idea of the development of the Ohře rift (Kopecký L. 1987), the Krušné hory fault played a subordinate role in the initial and main volcanic phases, and its more significant revival did not occur until the end of the Neogene or the Quaternary (cf. Elznic et al. 2010).

The link between the origin of the fracture and the rift is also indicated by the fact that in the regional geological structure of the foundation it does not express itself as an important structure or interface; nor even causes significant geophysical manifestations (Malkovský, 1985).

Fig. 2. Schematic cross-sectional structural section of the middle part of the Eger Graben (taken from Kopecký L. 1987). 1A - Tertiary sediments, 1B - Tertiary volcanics, 2 - Upper Cretaceous sediments, 3 - Lower Paleozoic sediments, 4 - Lower Paleozoic volcanics, 5 - Metamorphic Upper Proterozoic formations of the Teplá-Barrandian block, 6 - Metamorphic formations of the Krušné hory mountains,

7 - mylonite zone, 8 - presumed fault of pre-Tertiary age, 9 - pre-Tertiary phenitization.

History of research: The name of the fault was introduced by Suess (1903). Label Ore Mountain Fault, used in connection with the storage conditions of Upper Cretaceous sediments by Zahálka (1914), did not take place. In connection with the volcanism of the Central Bohemian Highlands, Hibsch (1891) named the eastern part of the Krušné hory fault.

As part of the post-war massive increase in interest in geology, several stages of research took place, in which the construction of the North Bohemian basins, incl. Krušné hory fault was paid considerable attention to. Immediately after World War II, the stage of searching for, verifying, and extracting raw materials from own resources began due to the limited possibilities of importing raw materials from the surrounding, by war similarly damaged countries. During the 1960s, the state's raw materials policy stabilized and most of the emerging, largely archival research works / reports already included quality study results even of those geological aspects that more or less went beyond the extraction of raw materials. Z. Brus, V. Cílek, J. Kavka, N. Krutský, P. Šantrůček, J. Šindelář and others were connected with exploration works on raw materials.

Exploratory work on the exploration and verification of coal and other minerals was increasingly accompanied and supported by basic research in virtually all thematic areas related to the development and disturbance of the coal basins of northern Bohemia. Although they were primarily focused on the North Bohemian and Sokolov basins, most of them more or less touched on some aspect of the Krušné Hory Mountains fault. The authors of major works include, for example, Č. Bůžek, J. Čadek, L. Domácí, A. Elznic, Z. Feyfar, V. Havlena, Z. Hokr, S. Hurník, G. Kačura, E. Knobloch, M. Kužvart. ,, M. Malkovský, V. Škvor, M. Štemprok, J. Tyráček, J. Václ, J. Vachtl, M. Vilímek, O. Zelenka and others. M. Váně represented an extraordinary expert on the issue of construction of the wider surroundings of the Krušné hory fault as a representative of industrial geology, who was able to keep up with the research sphere in terms of the quality of work and at the same time excelled in unrivaled local knowledge and amount of works.

After J. E. Hibsch, L. Kopecký comprehensively dealt with the issue of the development of the Eger Graben and especially the volcanism connected to it, followed by O. Shrbený and J. Ulrych. A. Kopecký and P. Vyskočil dealt with the issue of neotectonic activity of the fault. D. Marek, J. Rybář and B. Košťák, for example, dealt with engineering-geological issues associated with gravitational phenomena and the influence of mining on the slope of the Krušné Hory Mountains. Basic and deposit research of this period was accompanied by several stages of geological mapping of the state territory. The result of these works was a mapping of the entire course of the fault, finding out the basic features of both the fault itself and the storage conditions of its surroundings and the basic characteristics of the phenomena that are related to it.

Approximately since the early 1990s, another stage of research has been underway, when reducing pressure on industrial aspects, knowledge of the basic features of the geological structure of the Krušné Hory Mountains and, last but not least, significant development of technical equipment and research methods have led to the creation of works, which are significantly specially focused to solve individual problems. Selected works of this group are cited below. Deformation styles at the Krušné hory fault

According to Malkovský (1979), the typical features of the fault are the laterally changing character of the fault - flexure (Fig. 3) and also the changing direction.

Havlena (1982) and Hurník & Havlena (1984) raised a well-known and often published question concerning the nature of the Krušné hory fault. They, on the basis of the fracture-free removal of Lower Miocene sediments often observed in mining works, marked the Krušné hory fault as a great

fold (ie the character of flexure). According to this idea, the ascent of the Krušné Hory Mountains took place simultaneously with sedimentation in the Podkrušnohorské basins, accompanied by the faultless withdrawal of Miocene sediments. Only later did a system of faults emerge; the continuing uplift was realized by movements on them, accompanied by the squinting of the inclination of the drawn layers. The protruding layers reach vertical and even overturned inclinations in places. It is clear from the lithological characteristics of the seams that sediments could not have formed near the edge of the sedimentation pool, but originally continued through the Krušné hory fault to its northern surroundings. This refutes the assumption of some authors that the process of ejection took place in a synsedimentary manner with the formation of Lower Miocene formations (Elznic et al. 2010).

The dominant flexural character is mentioned by Malkovský (1977, 1979) in the section between Prunéřov and the former village of Kralupy near Chomutov. Here, according to the findings of Plzák (1968), the interface of Miocene sediments and crystalline rocks has the same slope as the slope of the Krušné Hory Mountains, on which its course was documented. The interface is interpreted as a continuous drawn-out base of sediments without breaking by a break at the foot of the slope. A similar continuous extension of the place to the vertical positions is documented in the next section of the break between Janov and the former. S. K. Neumann mine near Litvínov (Malkovský 1979).

However, in most of its course, the Krušné hory fault clearly has the character of a fault. Malkovský (1977) states that its western section, mostly in the Sokolov basin between Dolní Niva and Vernéřov, is typically a turning point. Similarly, the section in the North Bohemian Basin between Jirkov and Janov and further in the eastern part of the fault from the former S. K. Neumann mine near Litvínov via Osek, Střelná, Krupka to Jílové near Děčín again has the character of a fault (Malkovský 1979).

The second significant feature of the Krušné hory fault in the North Bohemian Basin are changes in its direction. Malkovský (1979) defines 4 basic sections according to the direction of the fault: between Litvínov and Střelná the fault has the direction NE - SW, between Střelná and Bohosudov the direction is E - W, between Bohosudov and Telnice the direction of NE - SW, and from Telnice to the east direction of E - W. It is noteworthy that the E - W direction acquires a fault where the subsoil is formed by granitoid rocks.

Architectural elements of the Krušné Hory Mountains Fault

More detailed characteristics of the breach in those parts of the fault where the fracture-free ejection is dominated are contained mainly in unpublished mining documentation. A published case is, for example, the work of Váně (1961), which describes the bending of a layer near Hrdlovka. The bend begins in the deepest part of the basin and from there in a section 1.5 km long, its slope gradually increases towards the Krušné hory slope up to 52 ° at the edge of the basin. In the bent part, the layer is significantly crushed (Malkovský 1985).

In parts of its course where the Krušné hory fault has a demonstrable character of dislocation, its characteristics were documented in several localities.

In the Sokolov basin, the fault has a predominant direction ENE - WSW and a steep inclination to S. The main fault and the zone of accompanying violations were uncovered in the former uranium quarry Odeř (Rojík et al. 2010). The exposed main fault had the character of a decrease with an inclination of 60 ° to S and an amplitude of about 300 m. Tertiary sediments were deformed into a flexure with an inclination of layers up to 60 °. However, in the fault zone there were also associated faults of the nature of thrusts. In the Sokolov basin, the fault is accompanied in some parts by basalt outflows (Hroznětín part of the basin), up to 20 m thick crushing zones and hydrothermal alterations with quartz veins (Rojík et al. 2010).

In the North Bohemian Basin, the characteristics of the main fault were directly observed and measured on several surface outcrops and technical outcrops of the fault and its zone of accompanying failure (Elznic 1963, Marek 1985, Váně 1985, Rybář 1997, Rybář & Košťák 1998, for more see chapter 7). In addition, archival reports mention a number of occasional and extinct outcrops, such as the outcrop at the housing estate in Střelná, borehole GÚ-35 (slope of the fault plane 40 °; to the south - see Malkovský 1985).

In the easternmost part around Libouchka, the Krušné hory fault has the character of a band up to 2 km wide, consisting of 3-4 sub-parallel faults (Adamovič & Coubal 2020).

The continuation of the fault towards the framework of the Krušné hory fault s.l. also has a character of a fault, where it passes into the Děčín fault field, formed by a set of faults in the direction ENE - WSW, limiting narrow parquet covers. The blocks are usually inclined at angles of 10 - 28 °, but they are documented up to inclinations exceeding 45 °. These faults are violated by a sub-equidistatic system of faults in the direction of the WNW - ESE, on which there were mainly right-hand sub-horizontal shifts, evidenced by striations.

In the section between the Děčín fault field and the Lusatian fault field (the so-called Doubice fault field), the Krušné hory fault s.l. has the form of one to two subparallel fractures with a relative decrease of SE blocks, which are in some places dislocated by faults of the E - W direction, the reactivation of which was significantly contributed by the horizontal shift (Adamovič & Coubal, 2020).

Fault structure and dip

…

Cross structures and Segmentation

…

Scarp morphology

…

Seismicity

To be revisited after completion of earthquake catalogue. Fix Me!

Pre-Miocene evolution

…

Fault activity in late Cenozoic

Tertiary

…

Quaternary

…

Related local evidence

*(See layer //**Local evidence*** on a map. The sites are listed in south-to-north order.)//

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Main data sources for fault map

…

Other notes

…

References

- Adamovič, J. & Coubal, M. (1999): Intrusive geometries and Cenozoic stress history of the

northern part of the Bohemian Massif. Geolines, 9, 5 – 14.
- Adamovič, J. & Coubal, M. (2020): Tektonická stavba. In Vařilová Z. (ed.): Geologie Českosaského Švýcarska. 109 – 131. Správa Národního parku České Švýcarsko. Ústí nad Labem.
- Bucha, V. (1993): Rozlišení zemětřesení a odpalů na lokálních seismických stanicích v severozápadních Čechách. PhD thesis, Přírodovědecká fakulta Univerzity Karlovy v Praze. 96 s. Praha.
- Coubal, M., & .Adamovič, J. (2000): Youngest tectonic activity of faults in the SW part of the Most Basin. Geolines, 10, 15 – 17.
- Coubal, M., .Adamovič, J. & Bendl, J. (1991): Kinematická, dynamická, genetická a časová analýza střezovského zlomu. MS Archív Čes. Geol. Služby. 66 s. Praha.
- Elznic, A. (1963): Severozápadní. omezení Chomutovsko - mostecko - teplické. pánve. Věst. Ústř. Úst. Geol., 38, 245 – 253.
- Elznic, A., Macůrek, V., Brož, B., Dašková, J., Fejfar, O., Krásný, J. Kvaček, Z., Mikuláš, R., Pešek, J., Spudil, J., Sýkorová, I., Teodoridis, V., Titl, F. (2010): Severočeská (mostecká) pánev. In: Pešek, J. (ed). Tertiary basins and lignite deposits of the Czech Republic. Czech Geological Survey, 40 – 137.
- Havlena V. (1982): Vznik podkrušnohorskych pánví. - Sbor. 8. exkurze uhel. geol. semináře, 5 – 9.Přírodovědecká fakulta UK a ČSMG Praha.
- Hibsch, J.E. (1891): Die Insel älteren Gebirges und ihre nächste Umgebung im Elbthale nördlich vonTetschen, Jb. K.-kön. geol. Reichsanst., 41, 235 – 288.
- Hurník S. - Havlena V. (1984): Podkrušnohorské hnědouhelné pánve a Krušné hory jako součásti neotektonické vrásové struktury. - Čas. min. geol., 29, 1, 55 - 67. Praha.
- Kopecký A. (1972a): Drobné vrásové deformace v nezpevněnych neogén-kvartérních sedimentech. - Sbor. geol. věd, G, 22, 117 - 151. Praha
- Kopecký A. (1972b): Hlavní rysy neotektoniky Československa. - Sbor. geol. věd, A, 6, 77 - 155. Praha.
- Kopecký L. (1987 - 88): Mlady vulkanismus Českého masivu. - Geologie a hydrometalurgie uranu,11, 3, 30 - 67; 11, 4, 3 - 44; 12, 1, 3 - 40; 12, 2, 3 - 56; 12, 3, 3 - 40. Vyzk. základna uranového průmyslu Stráž pod Ralskem.
- Král, V. (1968): Geomorfologie vrcholové oblasti Krušnych hor a problém paroviny. – Rozpravy ČSAV, řada matem.- přír. věd, 78, 9, 1 - 65. Praha.
- Král, V. (1985): Zarovnané povrchy České vysočiny. – Studie ČSAV, 10. 72 s. Academia. Praha.
- Malkovský, M. (1977): Důležité zlomy platformního pokryvu severní části Českého masívu. – Ústřední ústav geologický, 30 s.
- Malkovský, M. (1979): Tektogeneze platformního pokryvu Českého masívu. Knihovna ÚÚG, 53. 176 s. Praha.
- Malkovský M. (1980): Model of the origin of the Tertiary basins at the foot of the Krušné hory Mts.:volcano-tectonic subsidence. - Věst. ÚÚG, 55, 3, 141 - 150. Praha.
- Malkovský, M. et al. (1985): Geologie severočeské hnědouhelné pánve a jejího okolí. ÚÚG. 424 s. Academia, Praha.
- Malkovský, M. (1987): The Mesozoic and Tertiary basins of the Bohemian Massif and their evolution. Tectonophysics, 137, 31 – 42.
- Marek, J. (1985): Existuje krušnohorský zlom? Čas. Min. Geol., 30, 39 – 51.
- Plzák, V. (1968): Geologické poměry v prunéřovském výběžku chomutovsko-ústecké pánve. MS Geofond. Praha.
- Rajchl, M., Uličný, D. & Mach, K. (2008): Interplay between tectonics and compaction in a rift-margin, lacustrine delta system: Miocene of the Eger Graben, Czech Republic. Sedimentology, 55, 1419 – 1447.
- Rajchl, M., Uličný, D., Grygar, R. & Mach, K. (2009): Evolution of basin architecture in an incipient continental rift: the Cenozoic Most Basin, Eger Graben (Central Europe). Basin Research 21, 269–294.

- Rojík, P. (2004): Tektonosedimentární vývoj sokolovské pánve a její interakce s územím Krušných hor. PhD thesis, Přírodovědecká fakulta Univerzity Karlovy v Praze. 227 s. Praha.
- Rojík, P., Fejfar, O., Dašková, J., Kvaček, Z., Pešek, J., Sýkorová, I., Teodoridis, V. (2010): Cheb basin. In: Pešek, J. (ed). Tertiary basins and lignite deposits of the Czech Republic. Czech Geological Survey, 143-161.
- Rybář, J. (1997): Interpretation of data about technogenic activity at the toe of Krušné Hory Mts. affecting endogenous and exogenous processes in the rock environment. Acta Montana IRSM AS CR. Series AB, 106, 4, 9 – 24.
- Rybář, J. & Košťák, B. (1998): Open pit mining under Krušné Hory high slopes. 8th International IAEG Congress. Proceedings, 3027 – 3034. Balkema, Rotterdam.
- Stettner, G. (1974): The Bavarian part of the Bohemian Massif. In Mahel, M. (ed): Tectonics of the Carpathian Balkan regions. Geol. Úst. D. Štúra. Bratislava.
- Suess, F. E. (1903): Bau und Bild der Böhmischen Masse. In: Diener, K., Hoernes, R., Suess, F. E., Uhlig, V.: Bau und Bild Österreichs. F. Tempky, G. Freytag. Wien. Lipzig.
- Váně, M. (1961): Příspěvek k litostratigrafické pozici salesijských křemenců v severočeské hnědouhelné pánvi. Čas. Min.Geol., 3, 346 – 355.
- Váně, M. (1985): Geologická stavba podkrušnohorského prolomu a jeho tektonogeneze. Sbor. geol.Věd, Geol., 40, 147 – 181.
- Váně, M. (1999): Geologie Lounska pro třetí tisíciletí. – 471 s. Nákladem vlastním, Chomutov.
- Vilímek, V. (1996): Interpretation of new geodetic measurements on the fault slope (in Czech). MS , Přírodovědecká fakulta Univerzity Karlovy v Praze. XX s. Praha.
- Zahálka, Č. (1914): Útvar křídový v Českém středohoří. I. Text. – 232 s. Nákladem vlastním, Roudnice nad Labem
- Zeman, J. (1988): Charakter neotektonické morfostruktury Krušných hor a model jejího vzniku. Věst. Ústř. Úst. Geol., 63, 333 – 342.

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